



# Ecohydrology of Vegetated Catchments Under Climate Change

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**ANU**  
THE AUSTRALIAN NATIONAL UNIVERSITY



# Catchment responses to climate change

( $E_a$  = actual evaporation and  $Q$  = stream flow)

	Energy-limited				Water-limited		
	$E_a$	$Q$	Vegetation cover		$E_a$	$Q$	Veg'n cover
			Tropics	High latitudes/ altitudes			
$\uparrow P$	$\uparrow$	$\uparrow$	$\sim$	$\downarrow?$	$\uparrow$	$\uparrow$	$\uparrow$
$\downarrow E_p$ (e.g., $\downarrow$ wind, $\downarrow R_n$ )	$\downarrow$	$\uparrow$	$\sim$	$\downarrow$	$\downarrow$	$\uparrow$	$\uparrow$
$\uparrow [\text{CO}_2]$	$\downarrow$	$\uparrow$	$\sim$	$\uparrow?$	$\sim$	$\sim$	$\uparrow$

Expect a differential response to based on catchment location

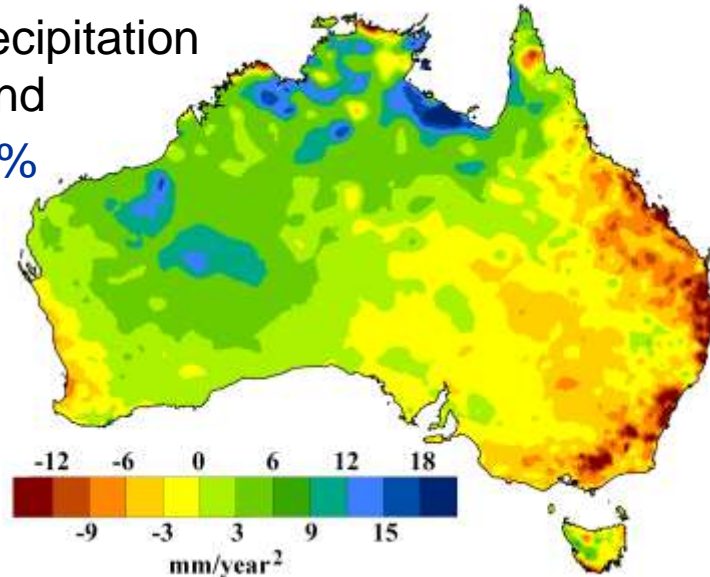
Use stream flow and remotely sensed vegetation cover (and  $E$ ) to assess if expectations are matched by observations

# Australia-wide trends in P and fPAR 1981-2006

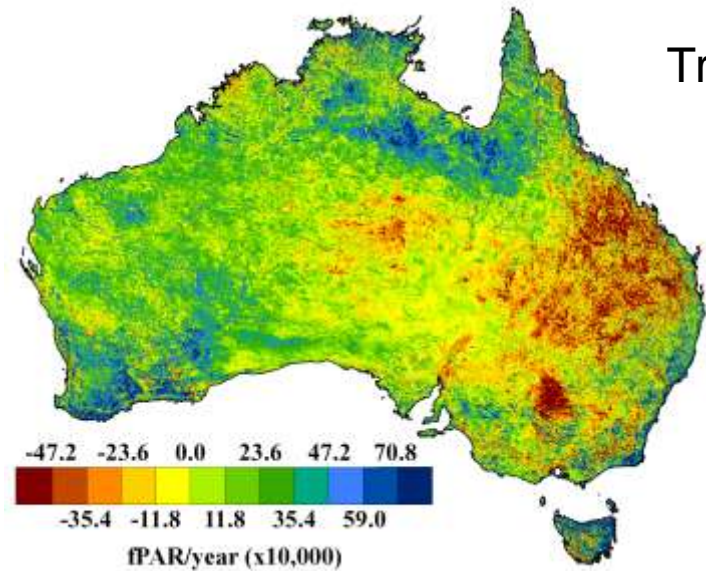
Donohue, McVicar and Roderick (2009) Global Change Biology

Precipitation  
trend

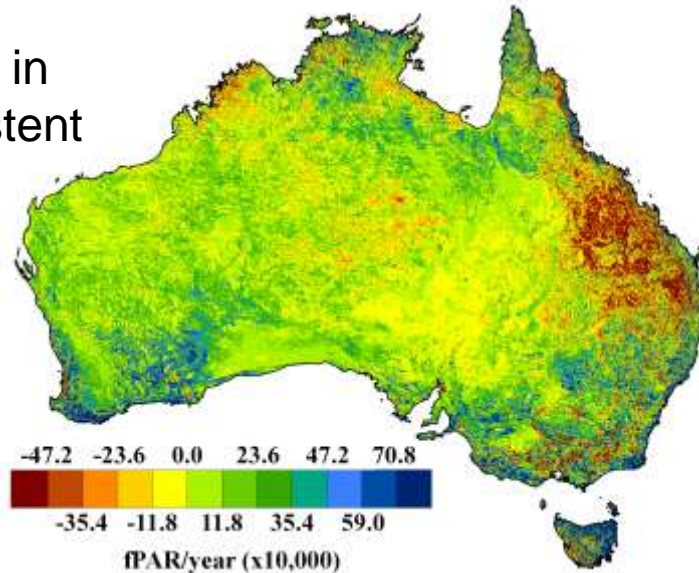
+7%



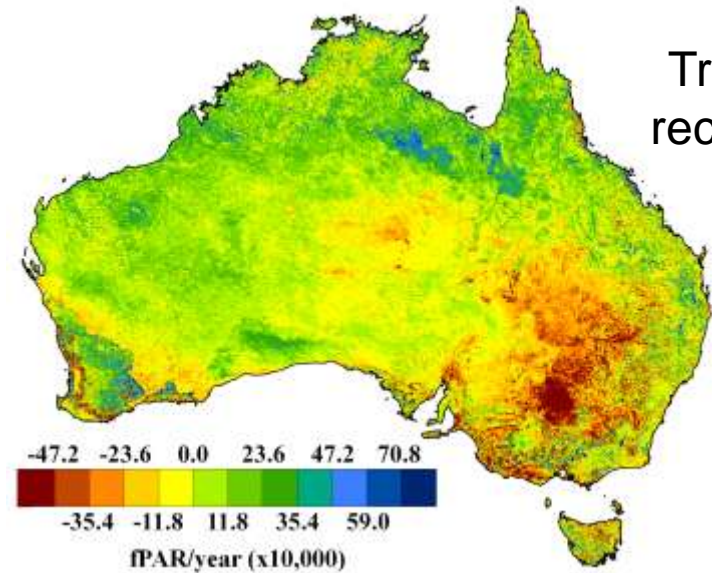
Trend in  
total  
fPAR  
+8%



Trend in  
persistent  
fPAR  
+21%

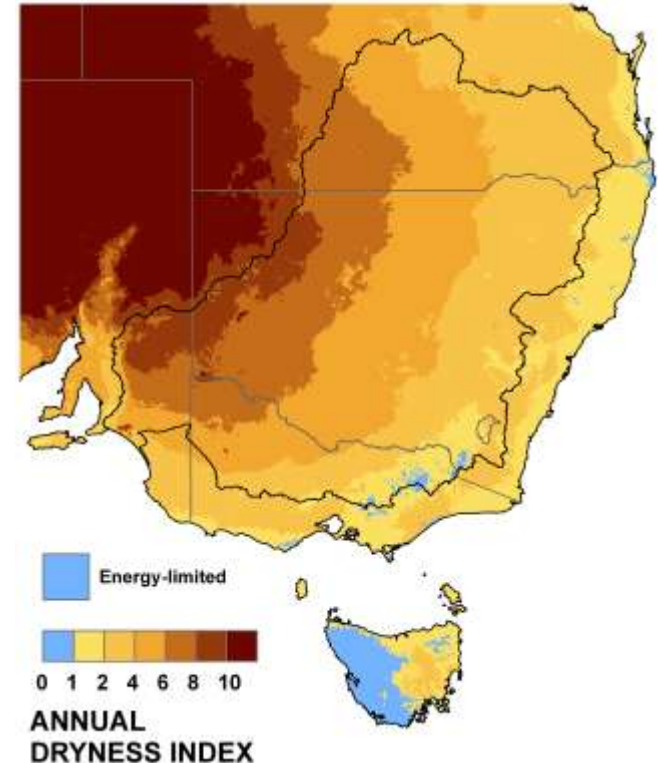
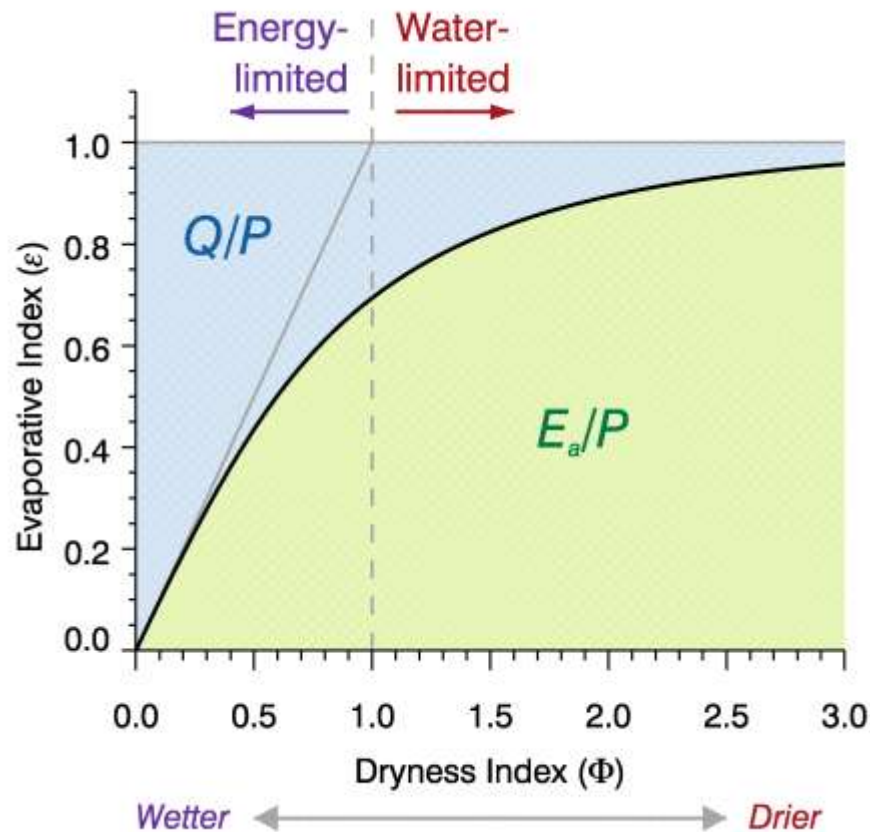


Trend in  
recurrent  
fPAR  
-7%





# Budyko's Framework : different locations, different responses

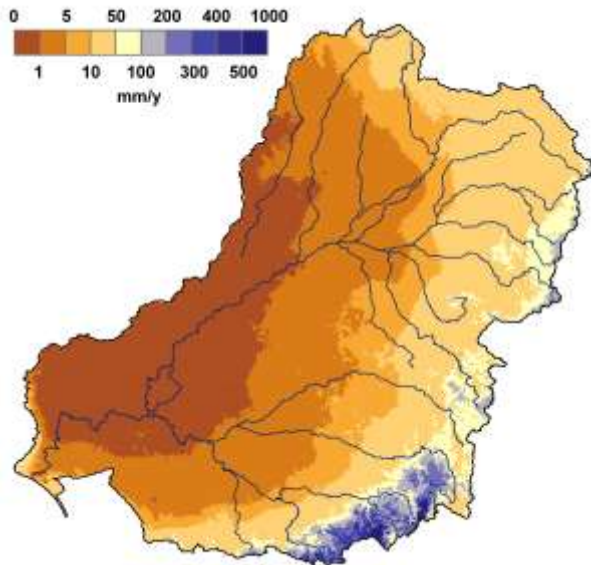


- The climatic Dryness Index ( $\Phi$ ) is the ratio of water demand to water supply;  $\Phi = E_p / P$
- The climatic Evaporative Index ( $\epsilon$ ) is the ratio actual evaporation to precipitation;  $\epsilon = E_a / P$

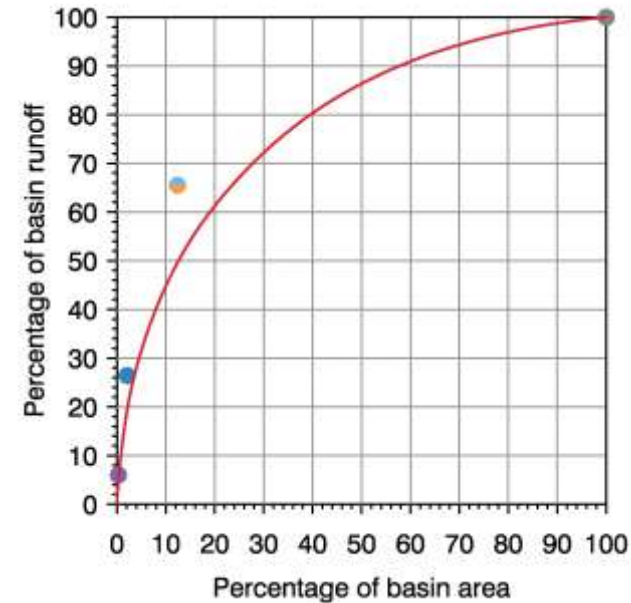
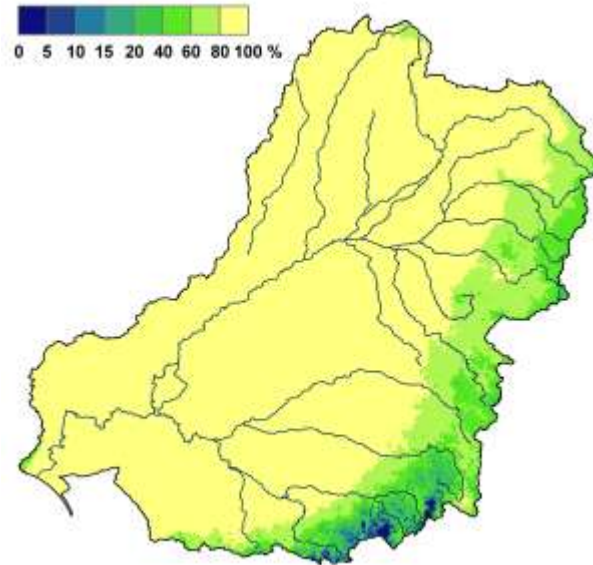
# Where does the Murray-Darling Basin's runoff originate?

Donohue, Roderick & McVicar (2011) J Hydrology – identify runoff producing areas

Annual Runoff



Runoff proportion by area



*R* is modelled using Budyko's curve and assumes no change in soil water storage

10% of the land area produces ~ 45% of total runoff

20% of the land area produces ~ 60% of total runoff

# Runoff sensitivity to CC using Budyko framework

Roderick and Farquhar (2011) Water Resources Research - developed method & applied in a lumped manner for the MDB. **An effective and transparent (+ repeatable) approach**

Choudhury formulation

$$R = P - \frac{PE_p}{P^n + E_p^n}^{1/n}$$

Essentially

$$R = f(P, E_p, n)$$

Sensitivity of Runoff

$$dR = \frac{\partial R}{\partial P} dP + \frac{\partial R}{\partial E_p} dE_p + \frac{\partial R}{\partial n} dn$$

Influence of 3 terms

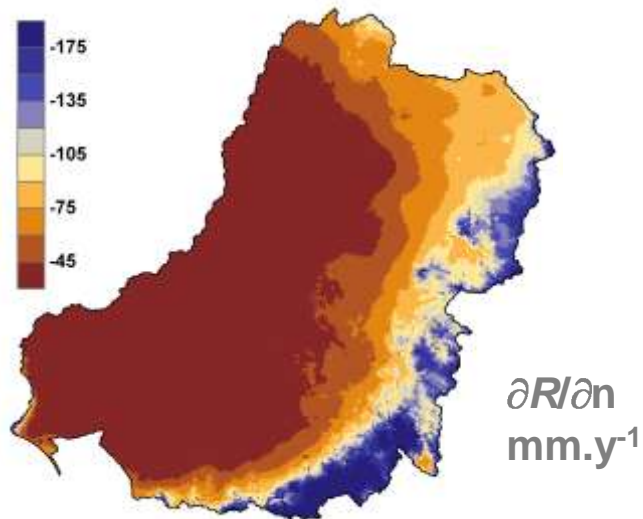
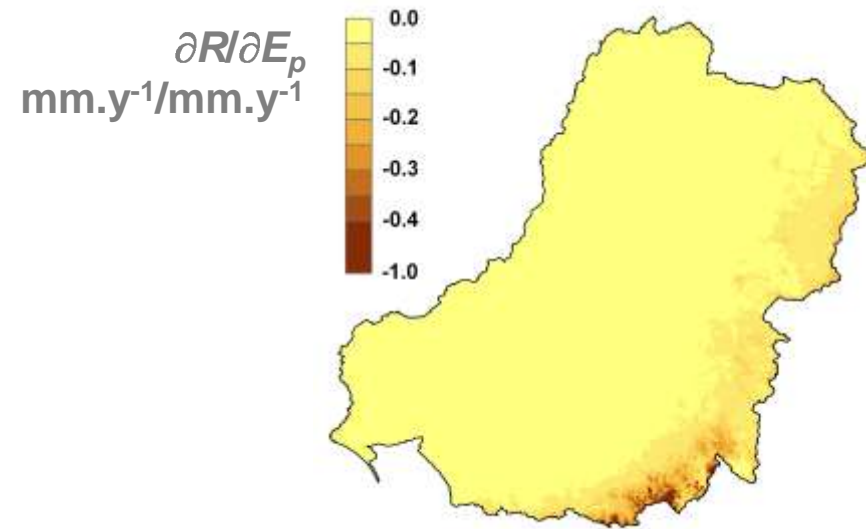
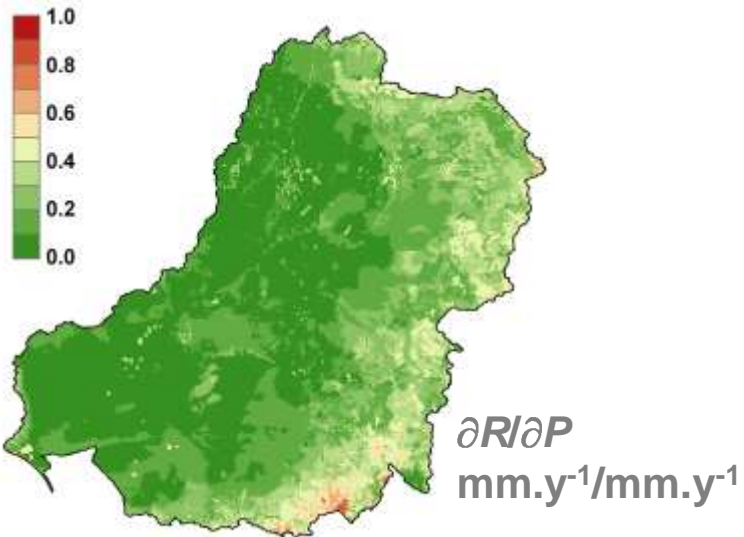
$$\frac{\partial R}{\partial P} = 1 - \frac{E}{P} \left( \frac{E_p^n}{P^n + E_p^n} \right)$$

$$\frac{\partial R}{\partial E_p} = \frac{-E}{E_p} \left( \frac{P^n}{P^n + E_p^n} \right)$$

$$\frac{\partial R}{\partial n} = \frac{-E}{n} \left( \frac{\ln(P^n + E_p^n)}{n} - \frac{P^n \ln P + E_p^n \ln E_p}{P^n + E_p^n} \right)$$

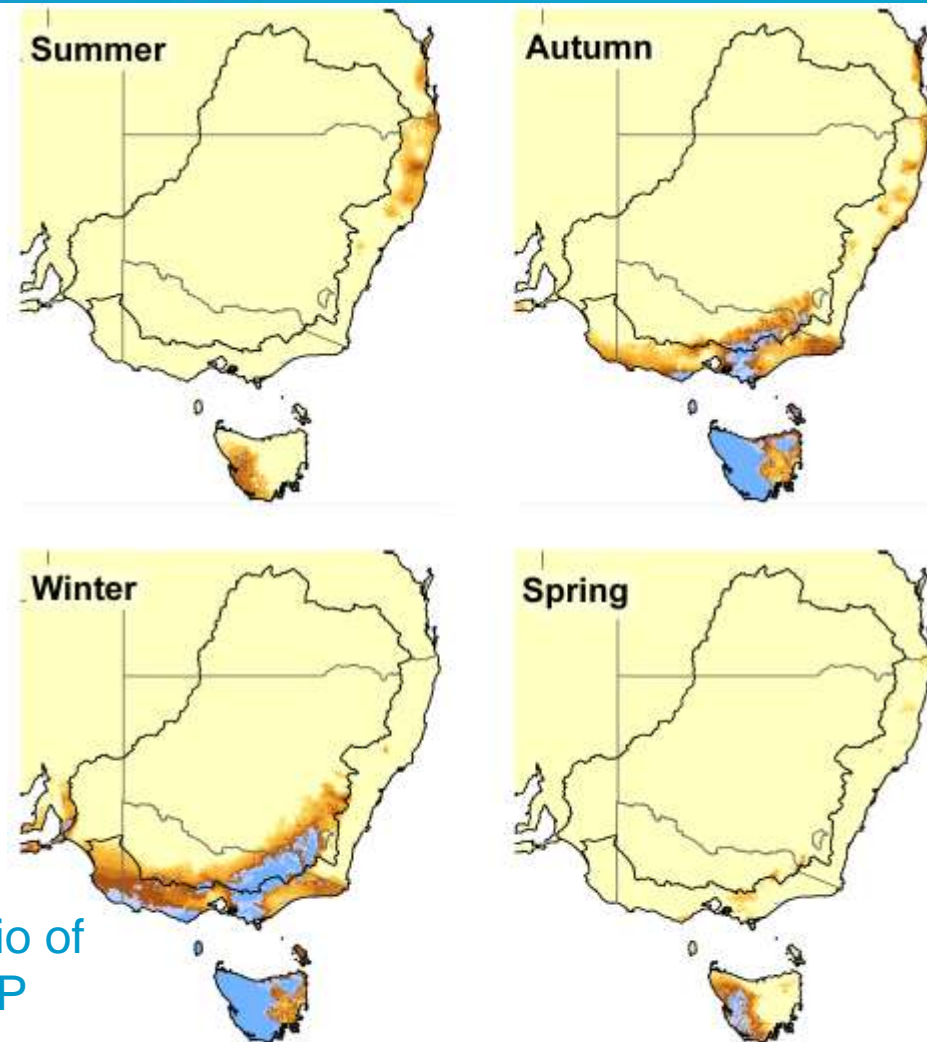
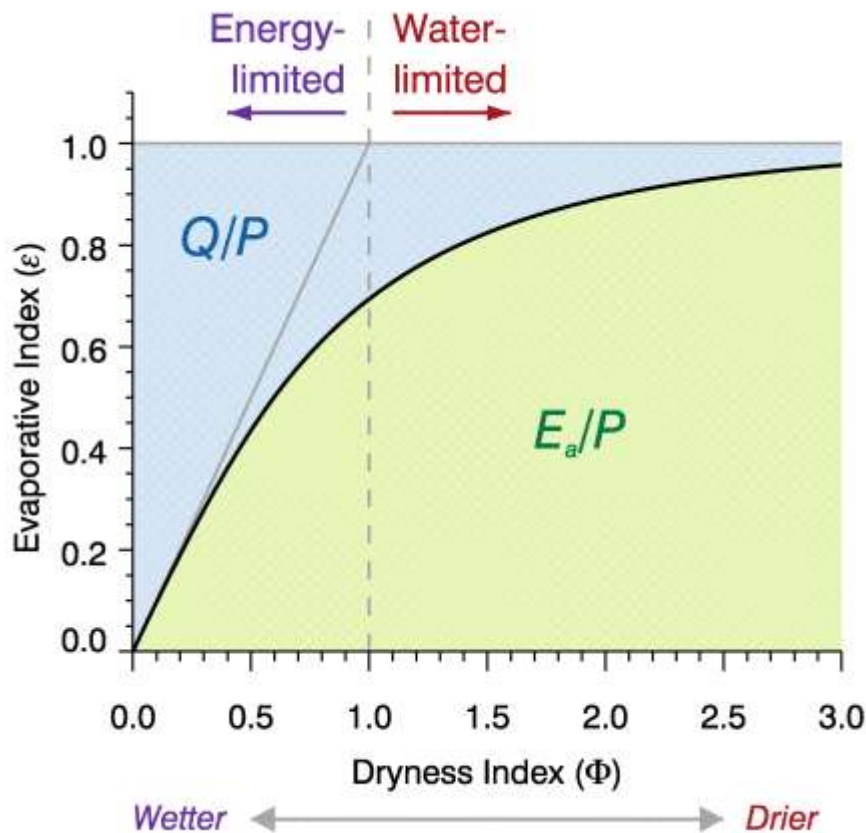
# Runoff sensitivity to CC using Budyko framework in the MDB

Donohue, Roderick & McVicar (2011) J Hydrology – spatially apply the sensitivity approach to Budyko framework the MDB

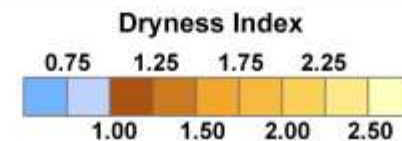




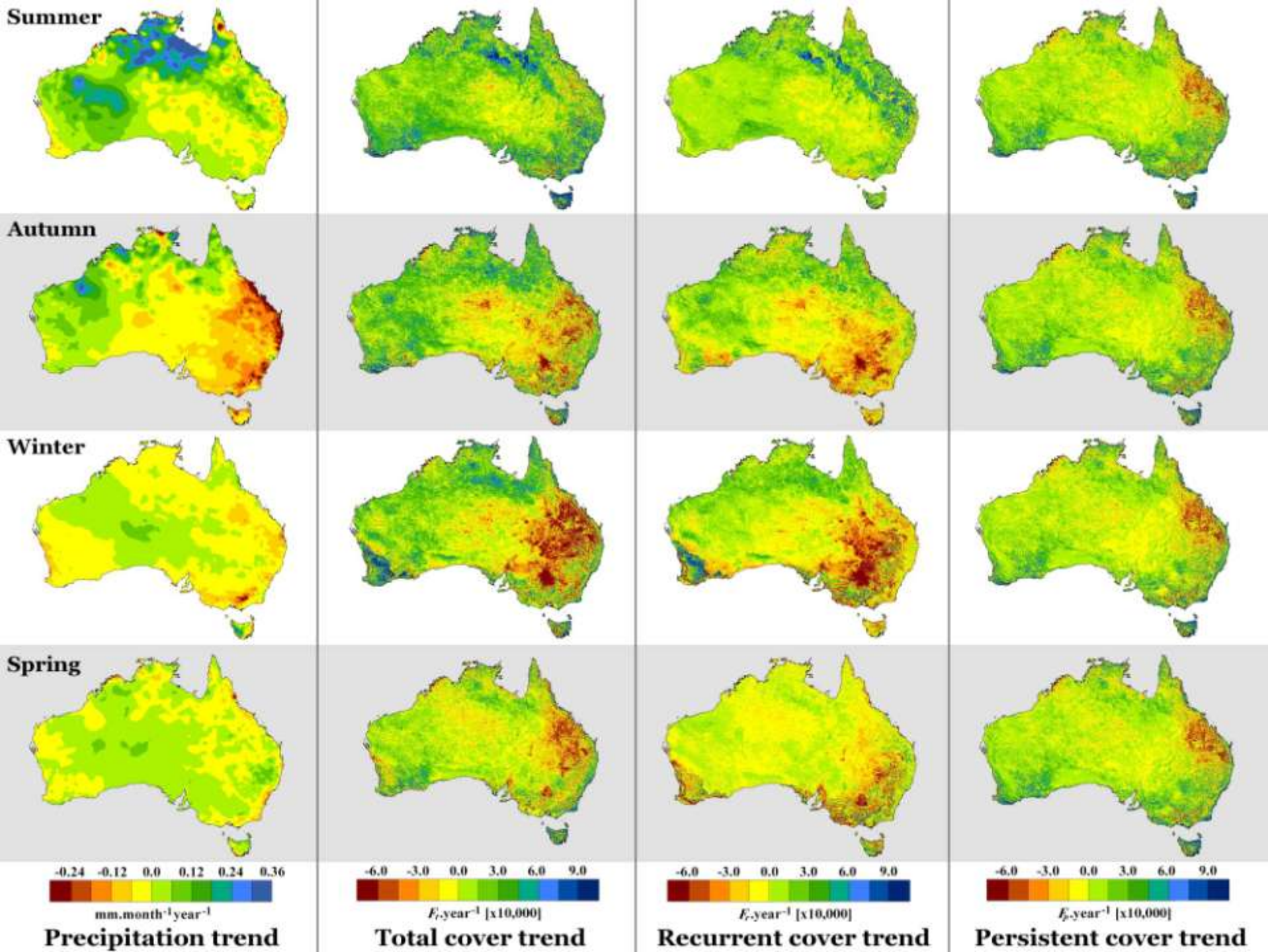
# Budyko's Framework : different locations + different times, leads to different responses



- The climatic Dryness Index ( $\Phi$ ) is the ratio of water demand to water supply;  $\Phi = E_p / P$
- The climatic Evaporative Index ( $\epsilon$ ) is the ratio actual evaporation to precipitation;  $\epsilon = E_a / P$



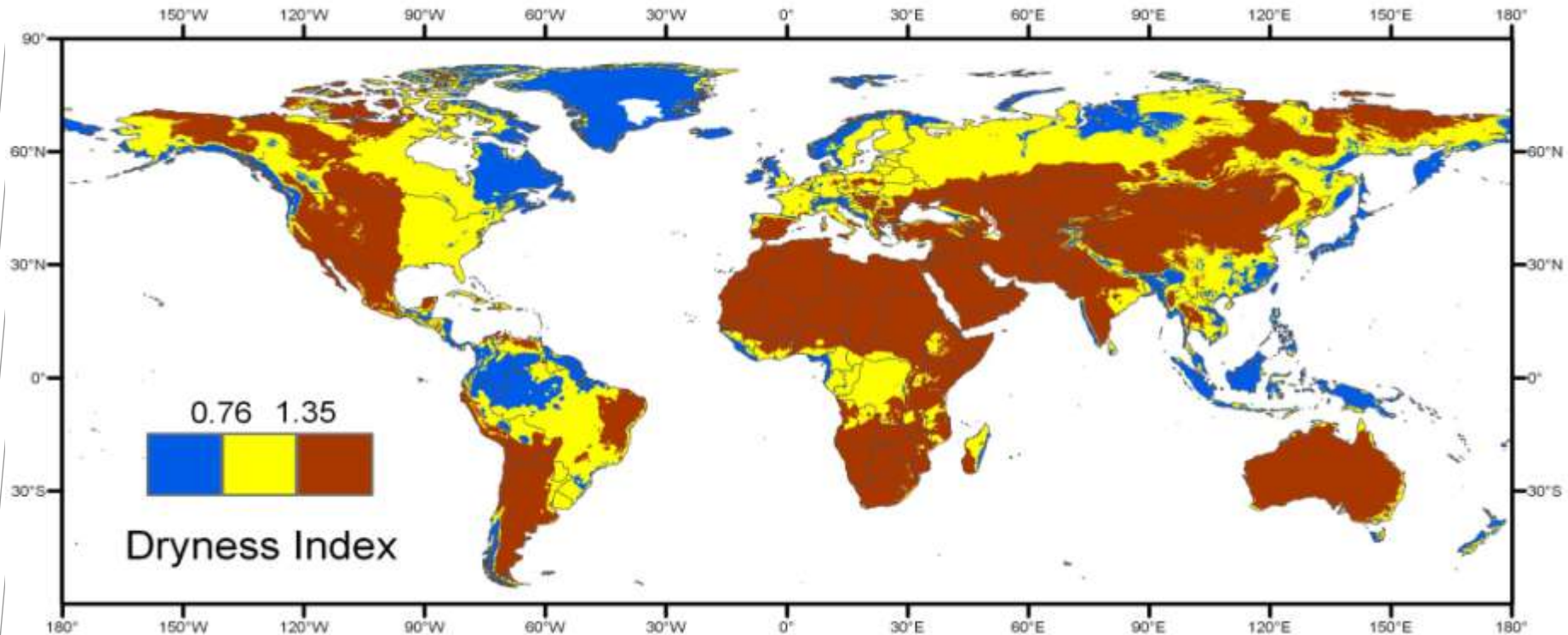
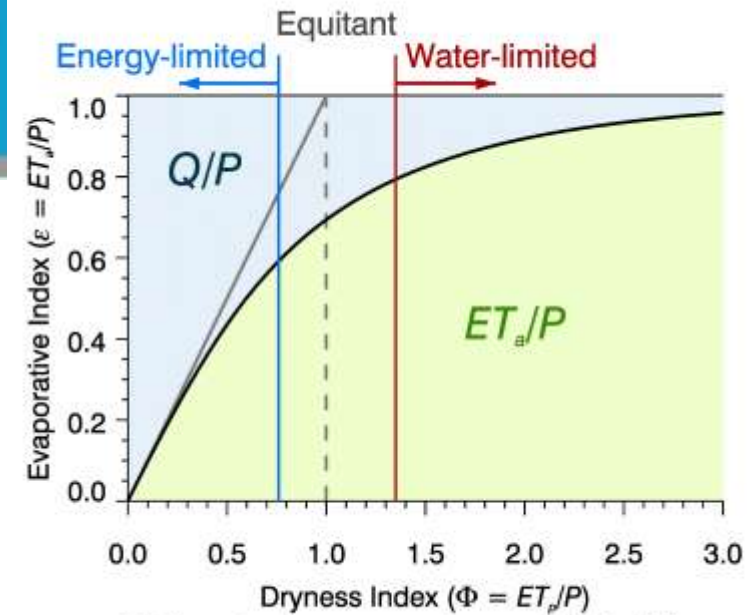




# Extending Budyko Framework: (Randall introduced BCP y/day)

Account for importance of seasonal dynamics:

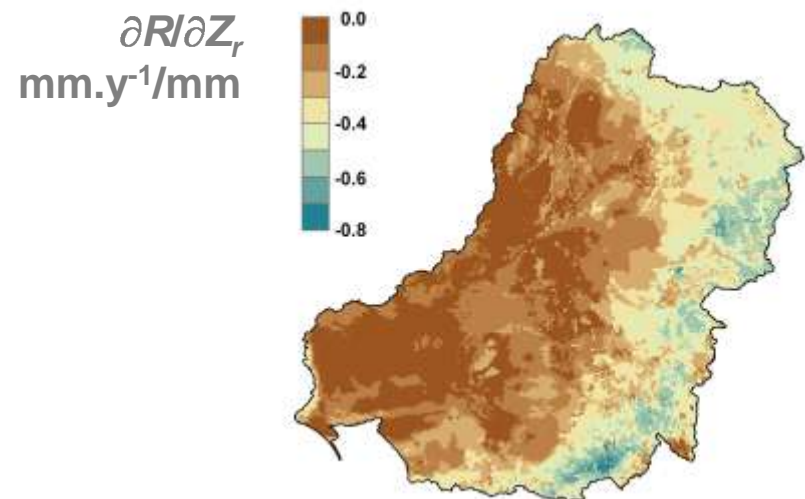
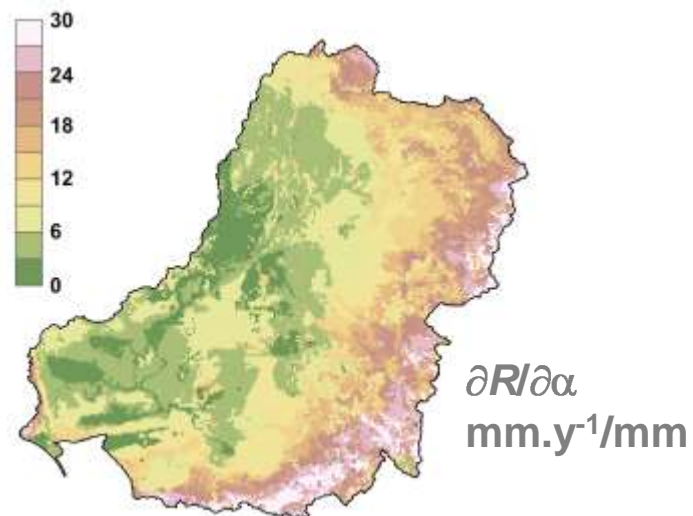
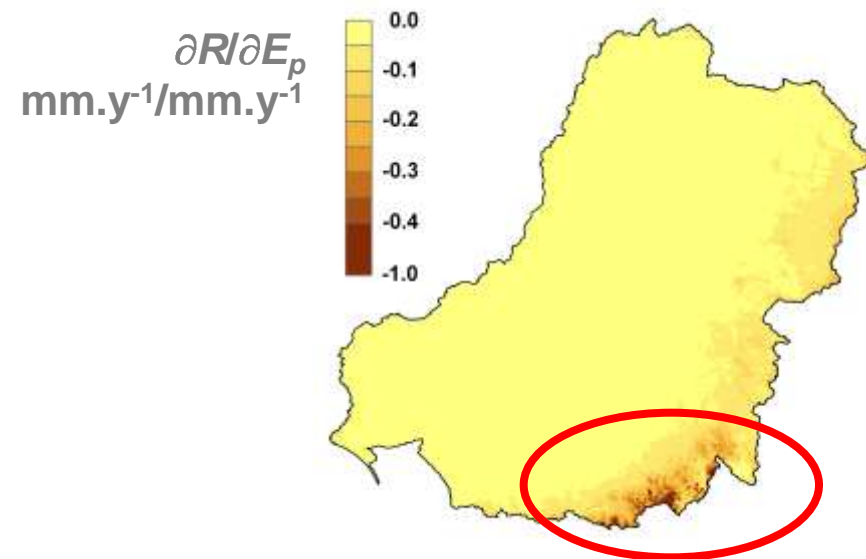
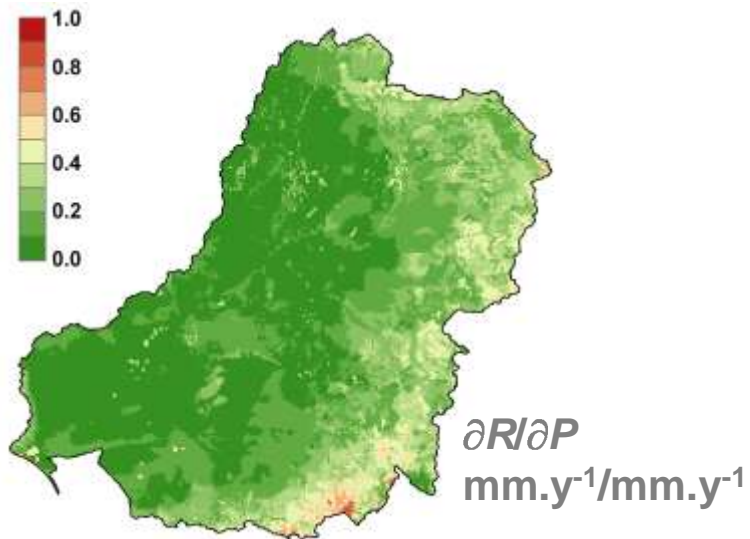
- 1) Precip and ETp
- 2) Vegetation activity (roots and phenology)
- 3) Storage (both soil and GW)





# Runoff sensitivity in the MDB using the BCP model

Donohue, Roderick & McVicar (2012) J Hydrology – developed BCP model (& inputs) and spatially assessed sensitivity of runoff to changes in inputs

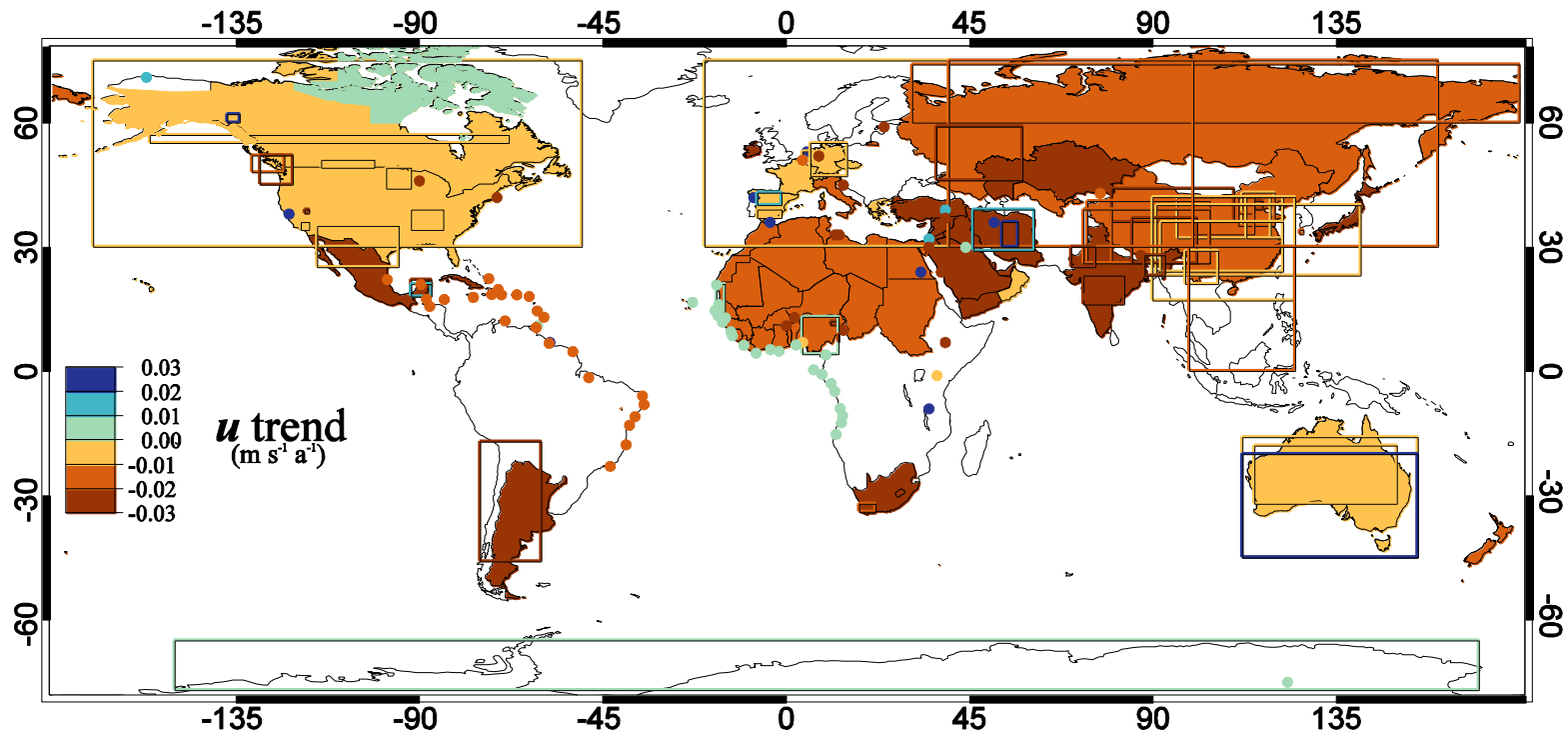




# Observed Global Trends in Near-Surface Wind Speed

McVicar, Roderick, Donohue et al. (2012) J of Hydrology

- Penman (1948) combined the radiative and aerodynamic components
- To date CC related evaporative trends had strong focus on Ta, Rad and Precip
- From 148 regional studies average wind speed trend =  $\sim -0.014 \text{ m s}^{-1} \text{ a}^{-1}$
- Assuming a linear trend this is a  $-0.7 \text{ m s}^{-1}$  decline over the last 50 years
- Trend of  $ET_a$  depends on limit to  $ET_a$  (P in WL areas, P and  $ET_p$  in EL areas)

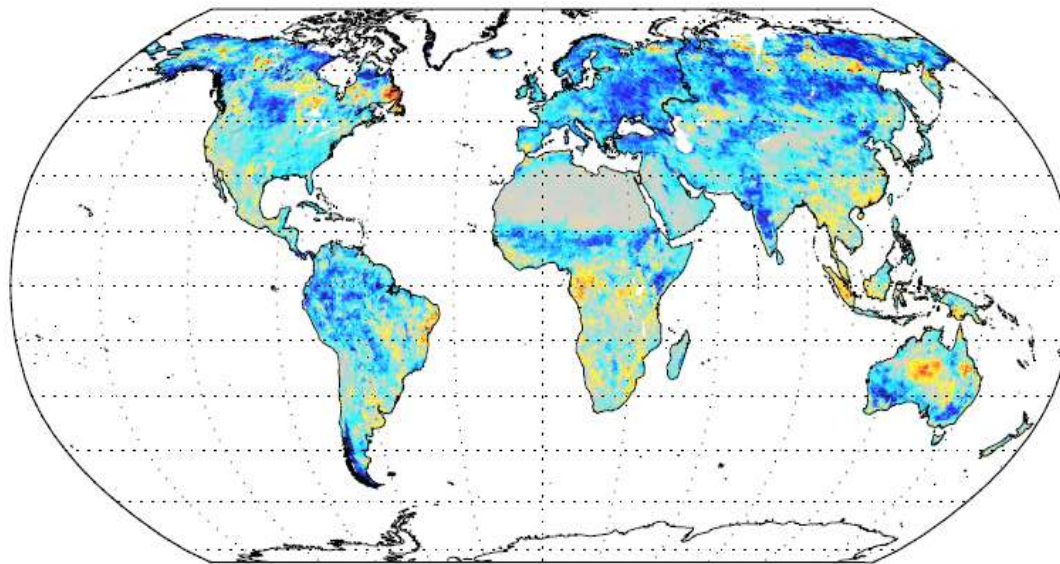


# Observed Global Trends in Near-Surface Wind Speed

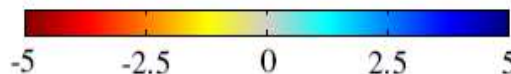
Most plausible drivers of near-surface wind speed trends are:

- i. increased roughness (Vautard et al., (2010) NGS);
- ii. widening of the Hadley cell (Seidel et al., (2010) NGS); and
- iii. increased aerosols (Jacobson and Kaufman (2006) GRL)

(b) GIMMS



Trend [ $10^{-3}$  NDVI units  $\text{yr}^{-1}$ ]



1982-1999  
AVHRR NDVI trends  
Beck et al (2011) RSE

GIMMS best able to track  
trends in 1424 Landsat  
pairs located worldwide

# Observed Global Trends in $E_{pan}$

McVicar, Roderick, Donohue et al., (2012) J of Hydrology

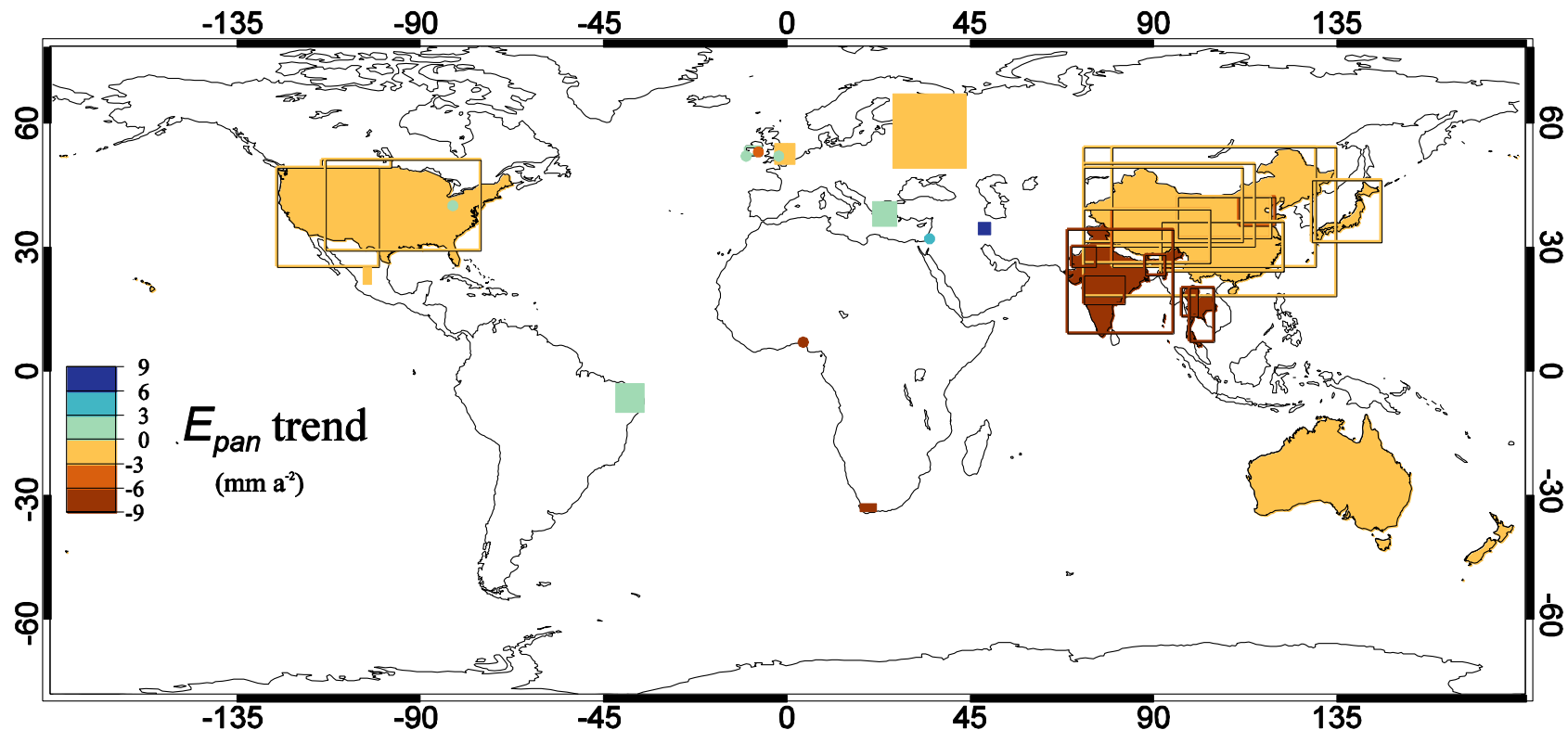


Globally 58  $E_{pan}$  studies reviewed with average trend =  $-3.2 \text{ mm a}^{-2}$

Only Class A pan studies =  $-3.8 \text{ mm a}^{-2}$  (n =37)

Studies using more than 10 Class A pans =  $-2.6 \text{ mm a}^{-2}$  (n =24)

Meteorological data forced modelled  $E_{pan}$  trends in line with observed  $E_{pan}$  trends in both N and S hemispheres; based on Roderick et al., (2007) GRL



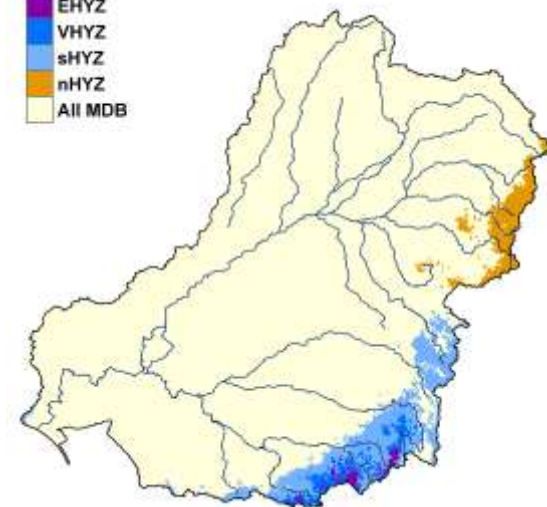
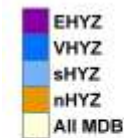


# Climate Non-Stationarity: $E_p$ Penman Trends (1981-2006)

Donohue, McVicar and Roderick (2010) J Hydrology

Area	Total (mm yr <sup>-1</sup> )	Trends of Penman potential evapotranspiration (mm yr <sup>-2</sup> )				
		Total	Due to $R_n$	Due to $u$	Due to $e_a$	Due to $T_a$
EHYZ	1248	-0.5	-0.3	-1.1	-0.5	+1.4
VHZ	1379	-0.1	-0.5	-0.7	-0.5	+1.6
sHZ	1502	-0.3	-0.6	-1.2	-0.4	+1.9
nHZ	1739	+2.0	+0.9	-0.7	-0.1	+1.9
MDB	1977	+0.4	-1.7	-0.9	+0.7	+2.3
All Aus	2340	-0.8	-0.6	-1.3	-0.4	+1.5

These southern water  
yielding zones provide  
50% of MDB Q



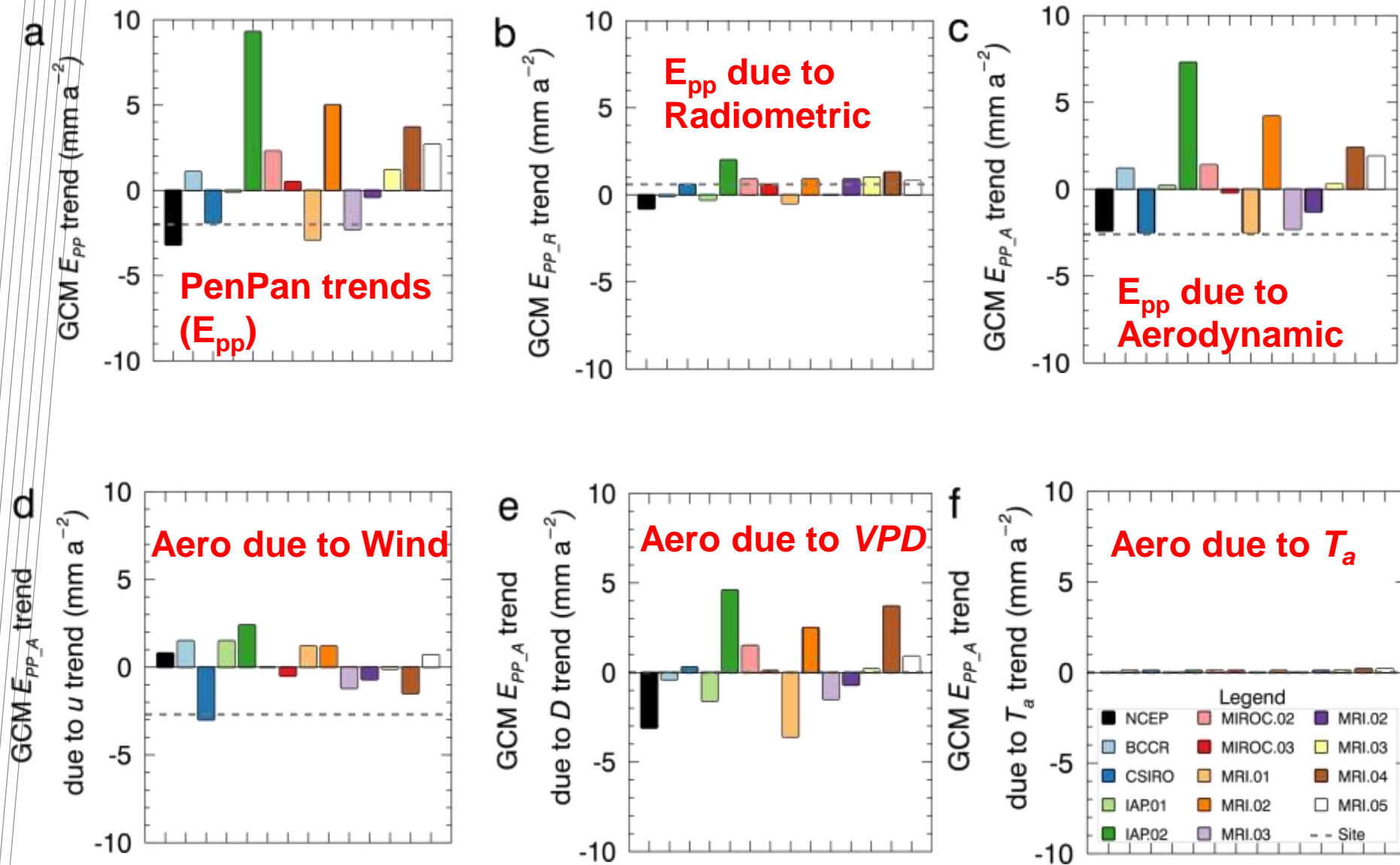
Different response to a changing climate depending on the location of the catchment / study area

There is not one rule for all locations for all time periods (responses change in space and time)

**Facing climate change not only global warming**

# Class A pan Evaporation : Attribution of the trends

Roderick et al (2007) GRL PenPan site data; Johnson and Sharma (2010) J HydroMet GCM data



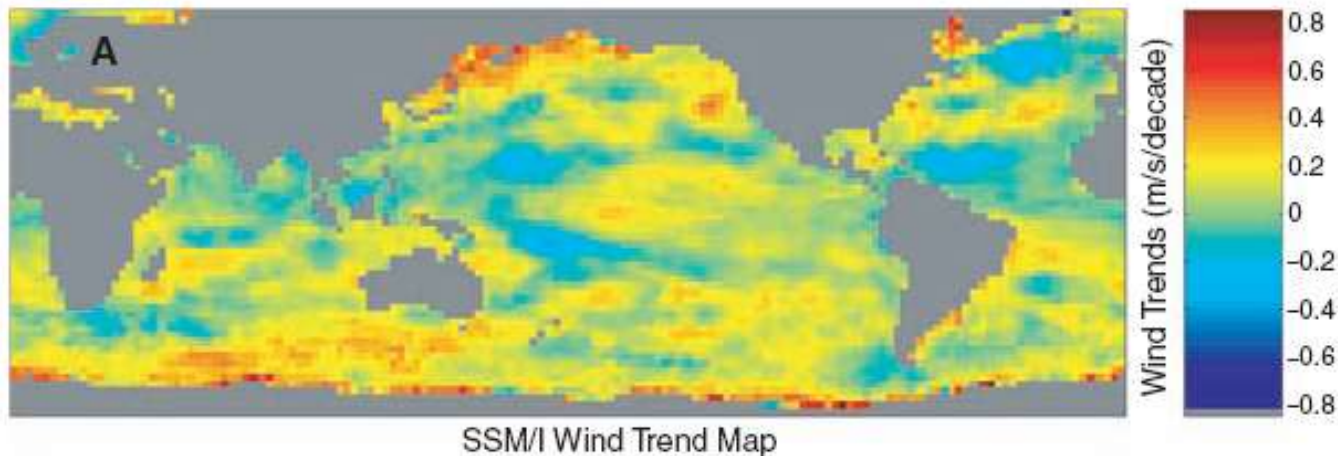
$T_a \uparrow$ ; Relative Humidity  $\sim$  Constant; Specific Humidity ( $q$ )  $\uparrow$

$q = g \text{ H}_2\text{O} / \text{kg wet air}$  and the mixing ratio  $= g \text{ H}_2\text{O} / \text{kg dry air}$

With much of the land-surface being EL, where is the water the coming from?

Intensification of the global hydrological cycle (Huntington 2006 J Hydrology)  
and are there greater rates of evaporation over oceans due to increased heat storage and wind speeds over the oceans?

Wentz et al (2007) Science: 1987-2006, average trend  $+ 0.008 \text{ m s}^{-1} \text{ a}^{-1}$





## CSIRO Land and Water

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[www.csiro.au/people/Tim.McVicar.html](http://www.csiro.au/people/Tim.McVicar.html)

Thank you